

# Propagating annular modes

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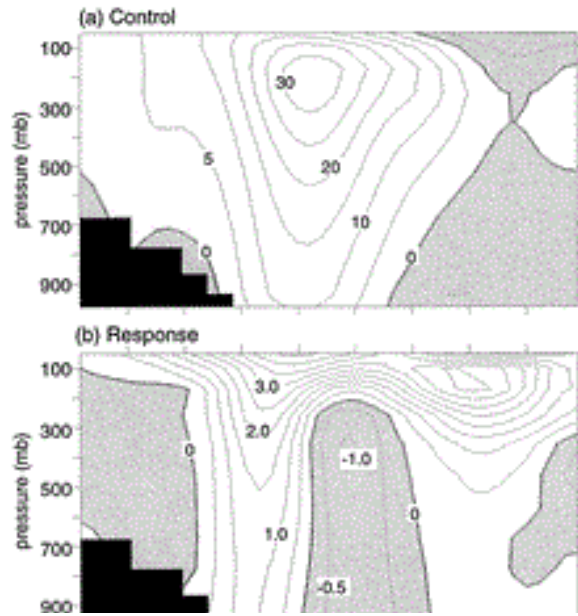
# Fluctuation-dissipation theorem

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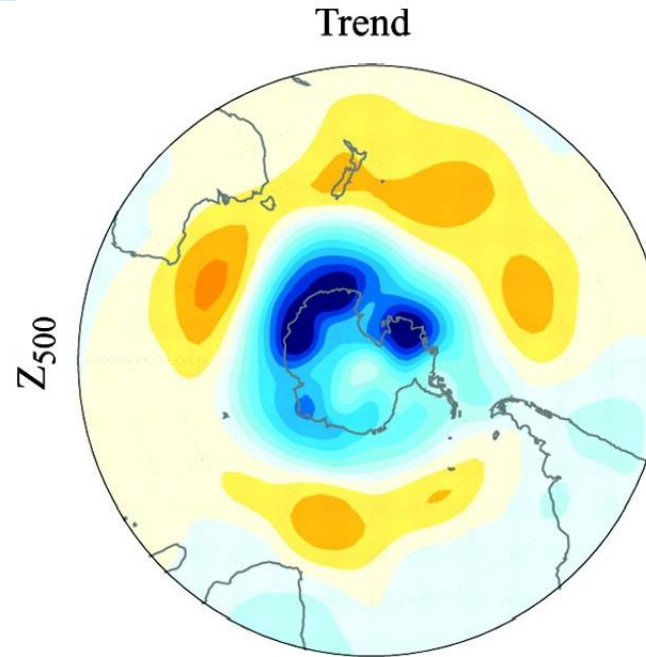
Relates the perturbed response of a dynamical system  
to its unperturbed variability

Response =  $\tau$  (mode timescale)  $\times$  projection of the forcing on the mode

# Dynamical climate responses and the fluctuation-dissipation theorem

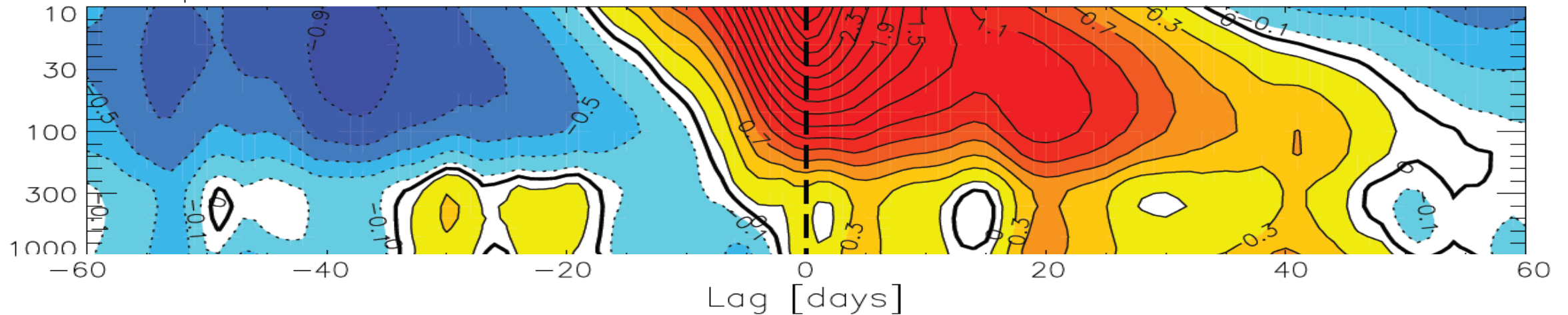


GHG  
Kushner et al. 2001



Ozone hole  
Thompson and Solomon 2002

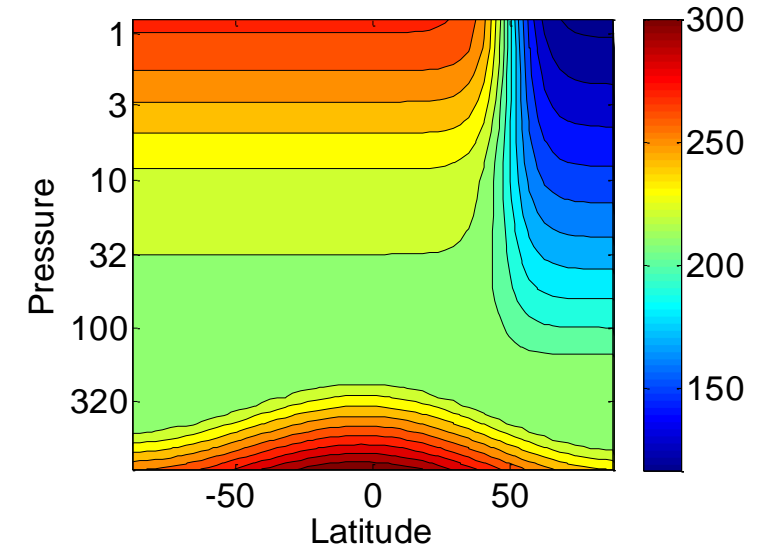
Composite of Warm Vortex Events



Kushner (2010)

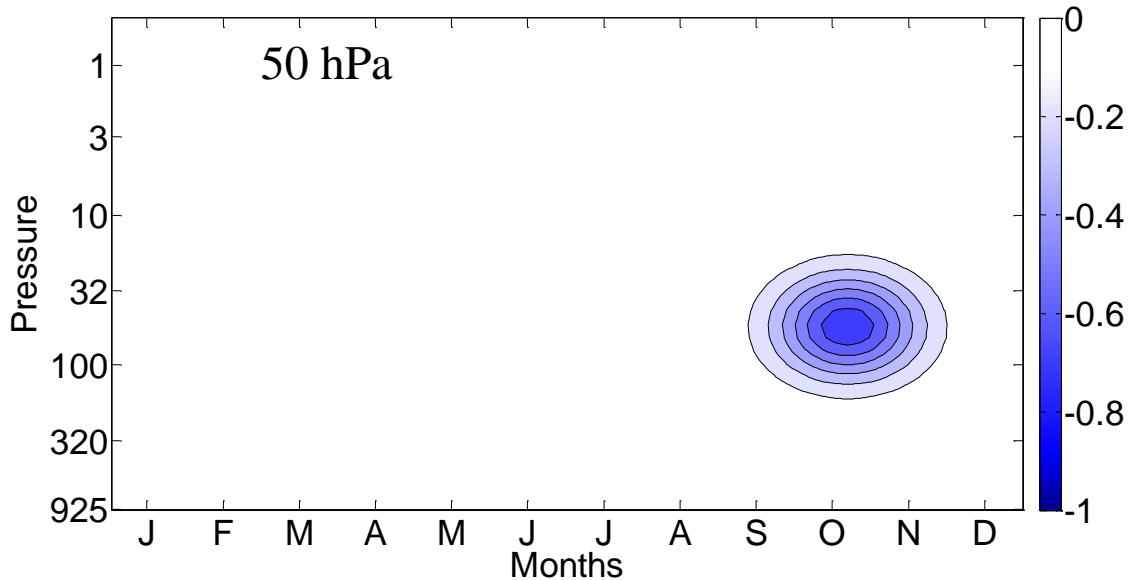
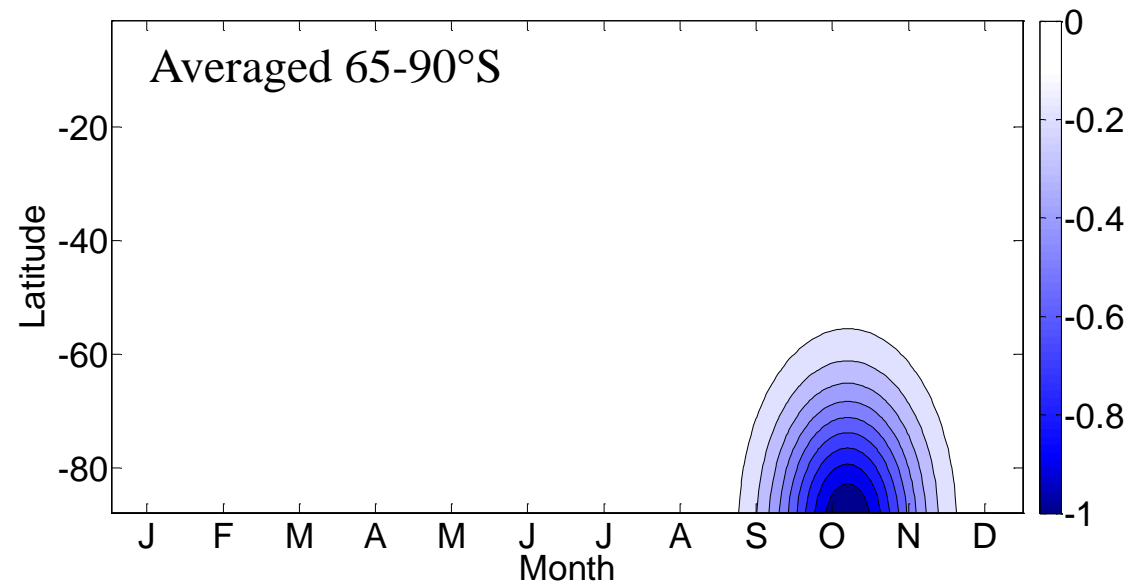
# Model setup

- Dry, hydrostatic, global primitive equation model (based on Polvani and Kushner, 2002), T42, 40 levels
- Replaces radiation and convection schemes with relaxation to a zonally symmetric  $T_{eq}$ , that is similar to Held Suarez in the troposphere
- In the stratosphere (above 200 hPa), equilibrium temperature set to U. S. Standard atmosphere, except over the winter pole, where a cold anomaly provides a representation of the polar vortex
- Simple seasonal cycle in equilibrium temperature that is **confined to the stratosphere**



$T_{eq}$  at NH winter solstice

# Polar stratospheric cooling

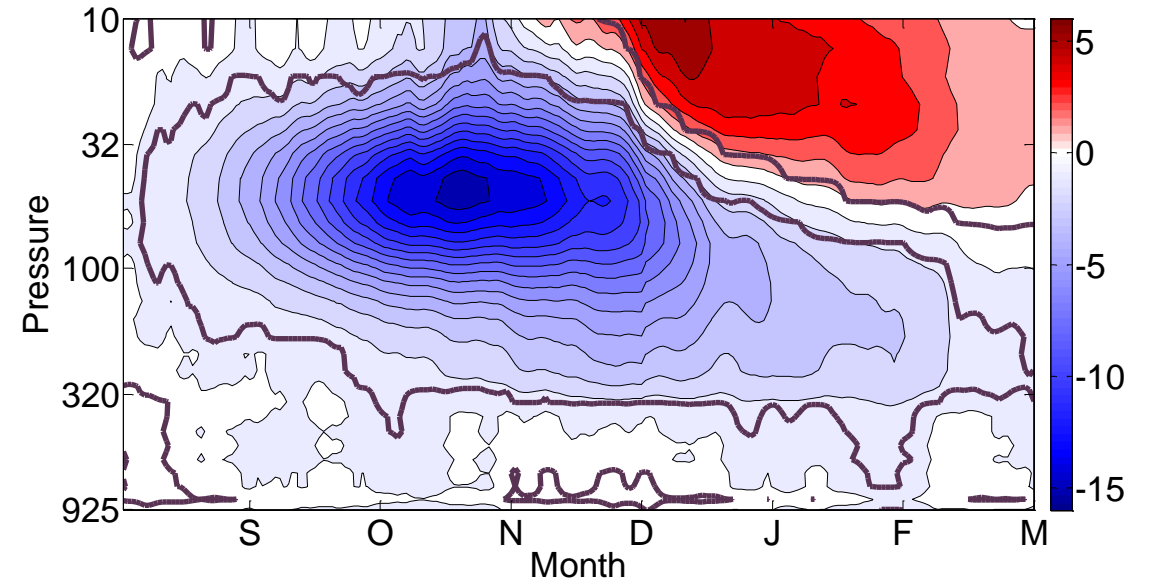


$$Q(\phi, \sigma, t) = q_0 \exp \left[ - \left\{ \frac{(\phi - \phi_0)^2}{2\sigma_\phi^2} + \frac{(-7000 \ln \sigma + 7000 \ln \sigma_0)^2}{2\sigma_\sigma^2} + \frac{(t - t_0)^2}{2\sigma_t^2} \right\} \right]$$

Similar to Butler et al. 2010 and Sun et al. 2014, but with

$$q_0 = -0.5K/day \quad \phi_0 = -\pi/2 \quad \sigma_\phi = 0.28 \quad \sigma_\sigma = 0.05$$

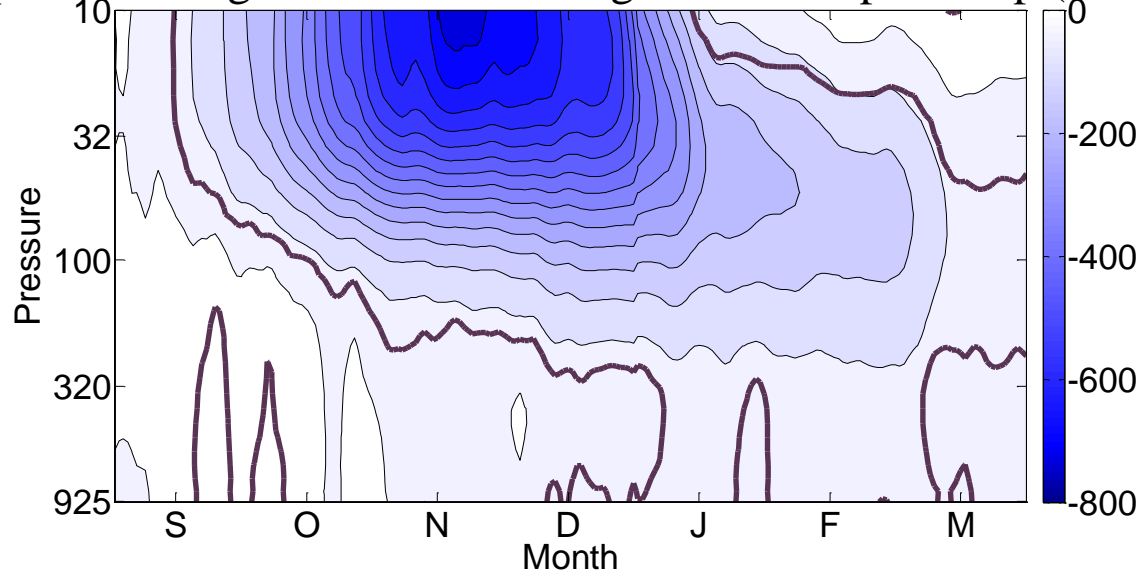
$$\sigma_\sigma = 4000 \quad \sigma_t = 20 \quad t_0 = \text{October 20} \quad (\text{We'll vary this})$$



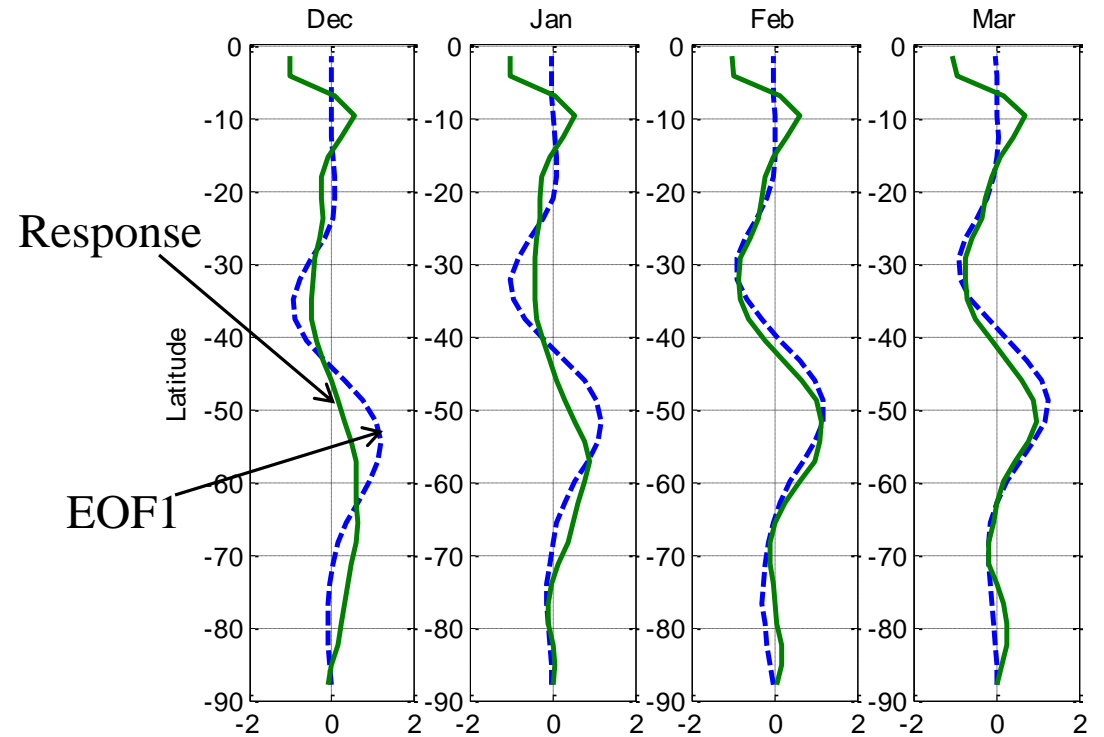
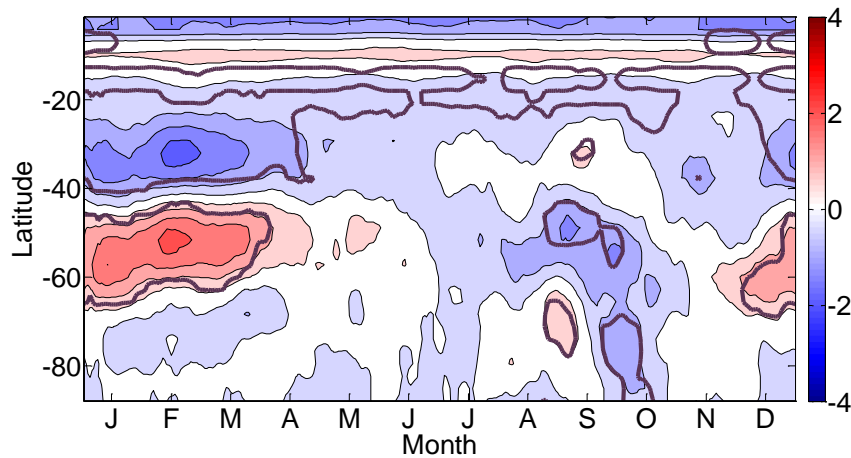
Temperature difference averaged over the polar cap (65-90°S)

# Tropospheric circulation changes

Geopotential height difference averaged over the polar cap (65-90°S)

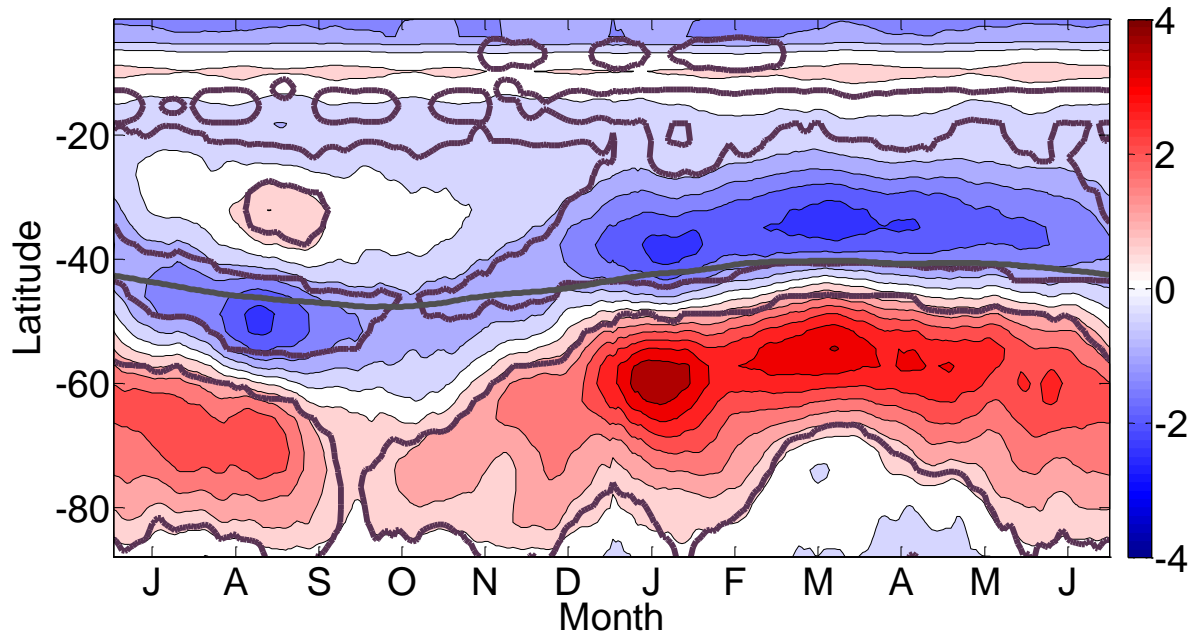


850 hPa [u] response

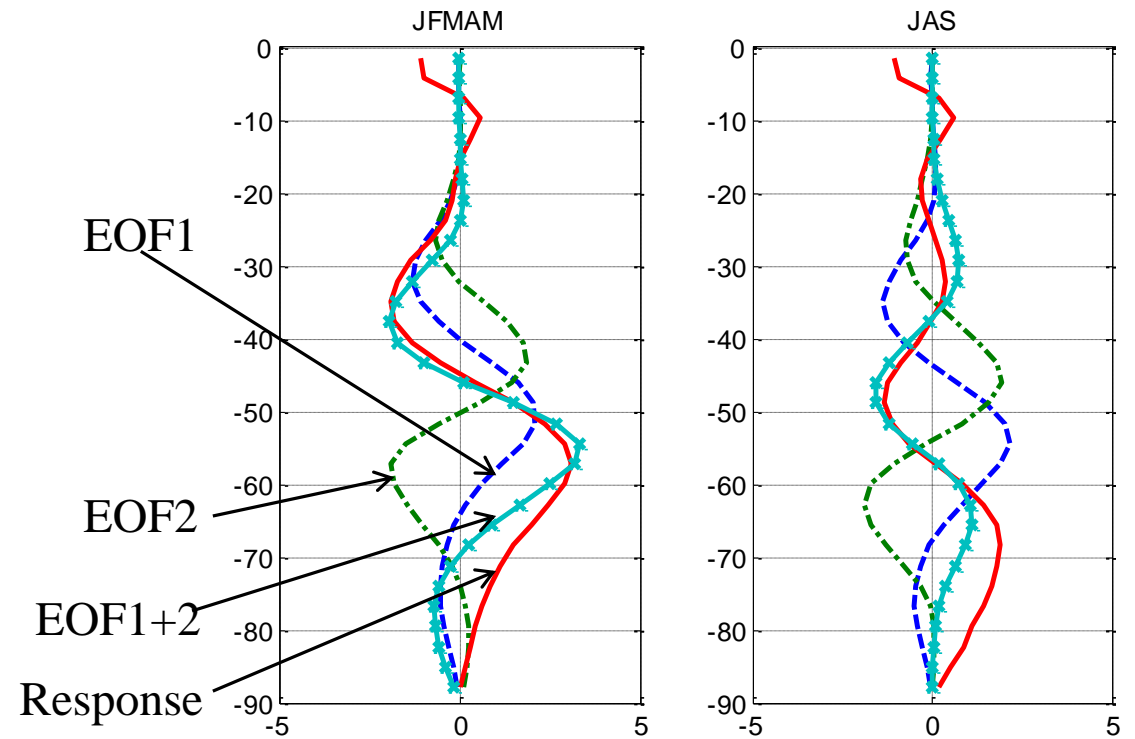


850 hPa [u] response doesn't always match the first annular mode!

# Surface responses to all year cooling – involves 2 EOFs

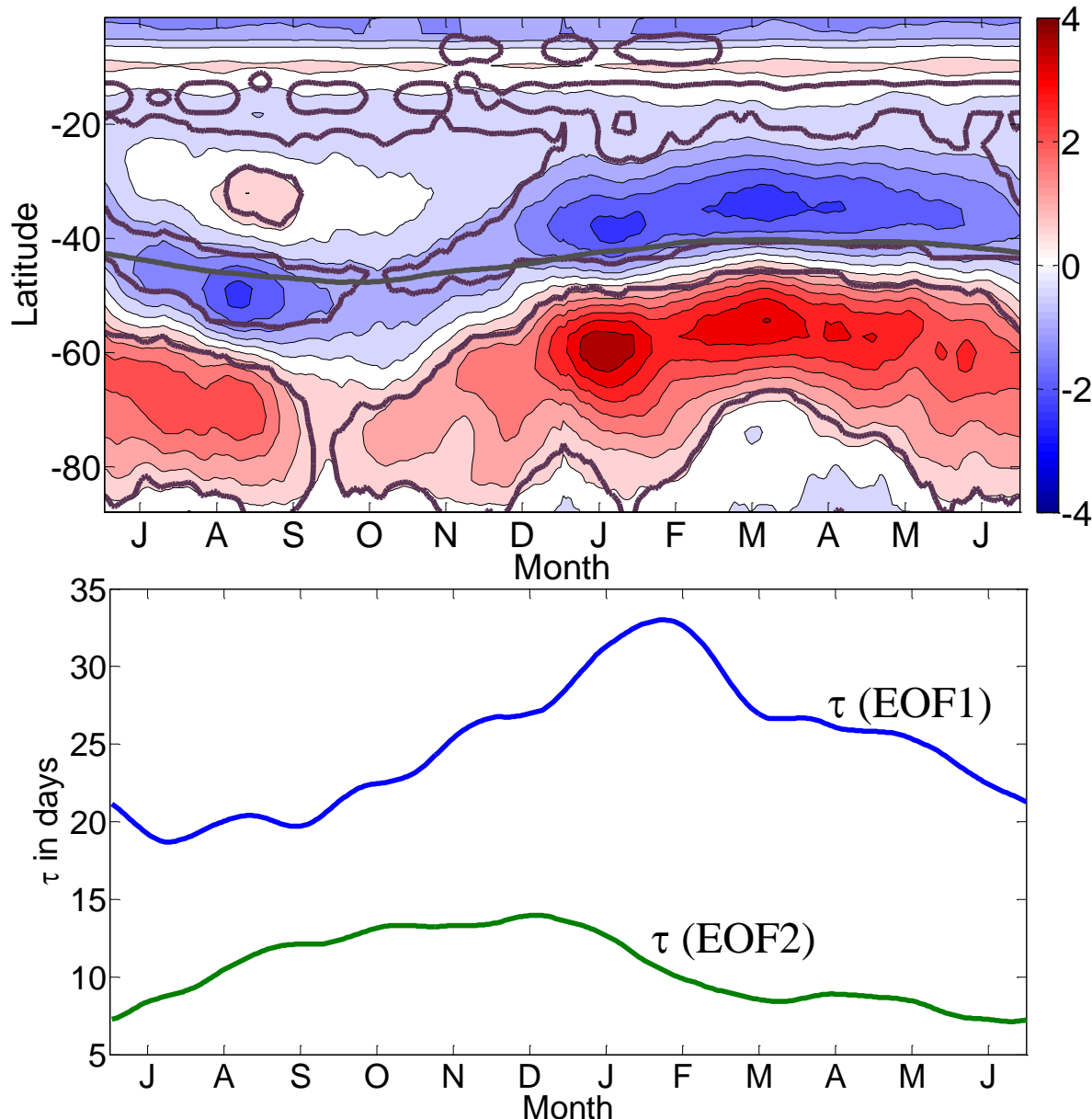


- Large dipole response early in the year
- Minimum, poleward shifted response in SON



Almost all the response explained  
by sum of contributions from 2  
EOFs

# Seasonal cycle of annular mode timescales

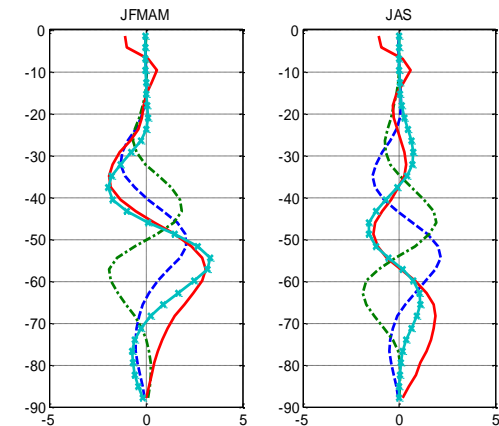
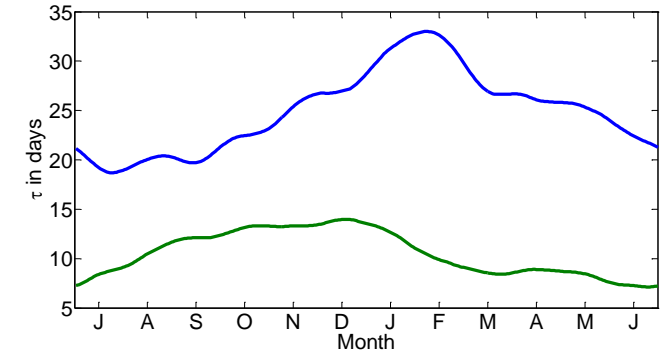


- Seasonality in tropospheric  $\tau$  (in the control run) induced by lower stratospheric seasonality
- If the projection of the forcing on the modes does not change through the year, the response in mode 2, relative to 1 would change by a factor of 2 over the year



# Summary – tropospheric response to ozone depletion

- Stratospheric seasonal cycle leads to seasonal changes in the persistence of the tropospheric mid-latitude jet and storm tracks
- The response of the atmosphere to external forcing cannot be described by a single “annular mode”!  
(cf. Black and McDaniel 2007a,b)



## Two leading EOFs actually one propagating mode?

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- **Principal Oscillation Patterns (POPs)** are the true modes

e.g. Von Storch (1988), Penland (1989)

$$G_\tau = C_\tau C_0^{-1}$$
$$G_\tau = V \Gamma_\tau W^T$$

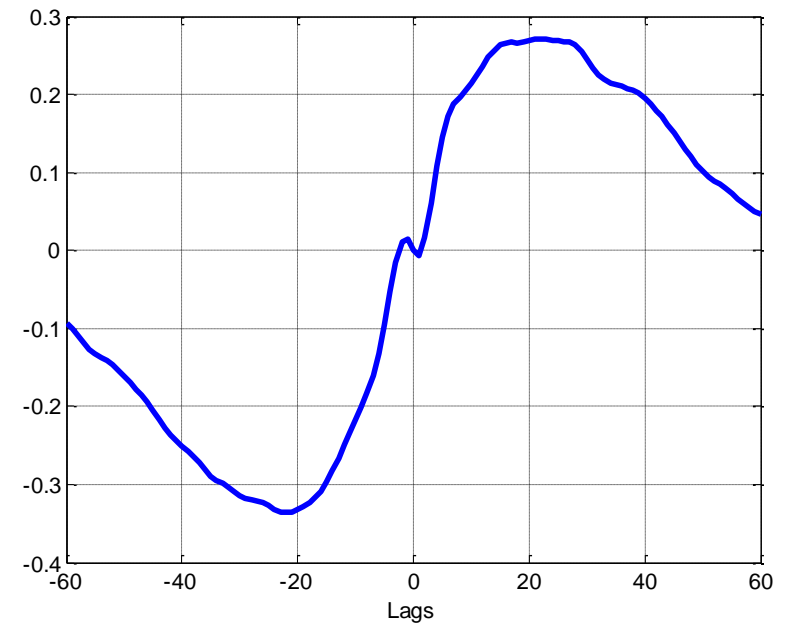
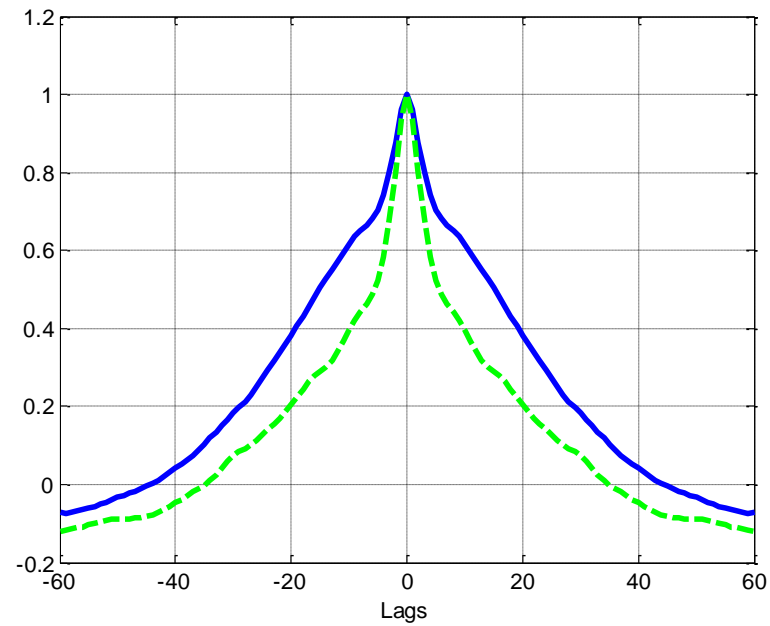
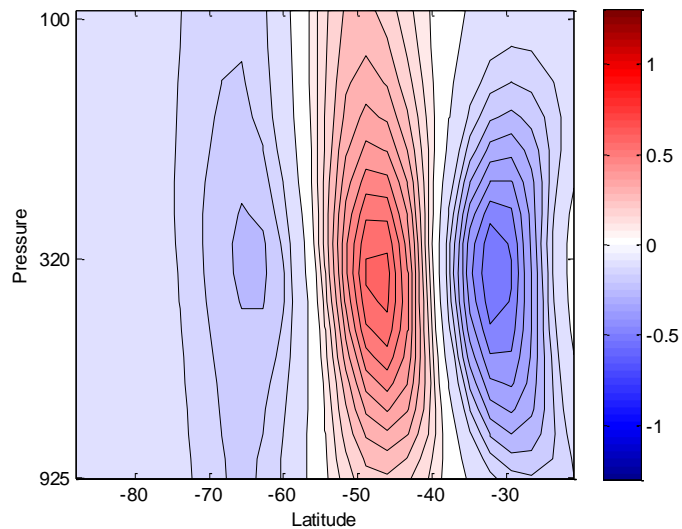
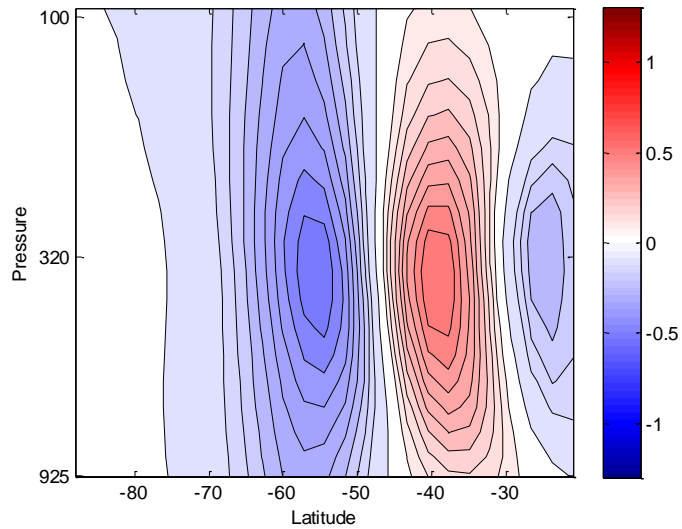
they contain dynamical information, and can be used to extract both **spatial patterns** and **timescales**. If EOFs were all independent, POPs would look like EOFs.

- Let's simplify the problem, and look at a perpetual winter situation

Are the model's **two leading EOFs** actually **one mode** describing **systematic latitudinal migrations** of the jet?

# PCs of two leading EOFs show lagged cross-correlation

2 leading EOFs explain 40% and 29% of variance, with  $\tau = 19$  and 13 days

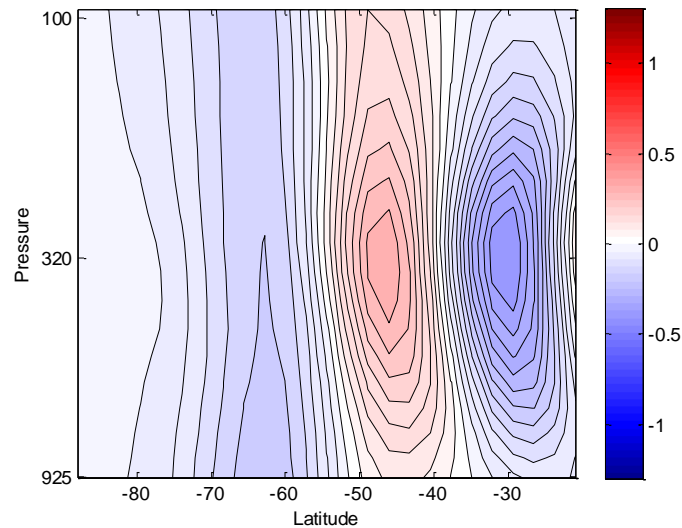
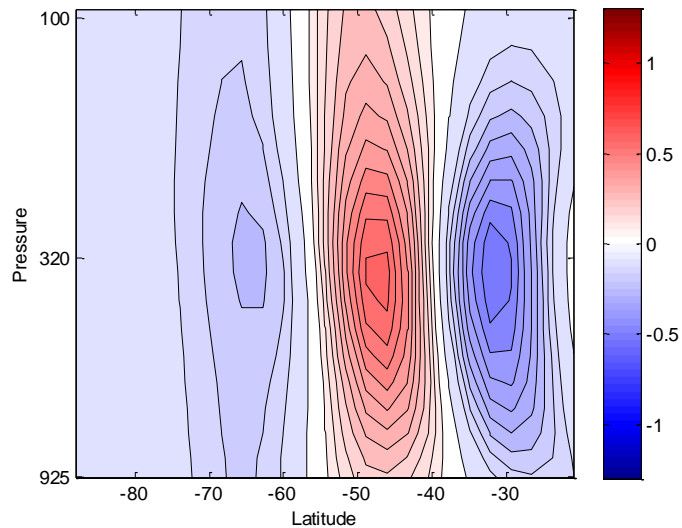
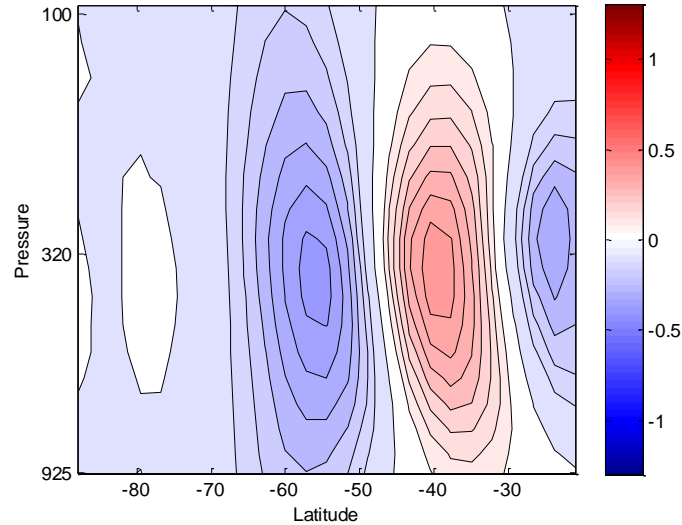
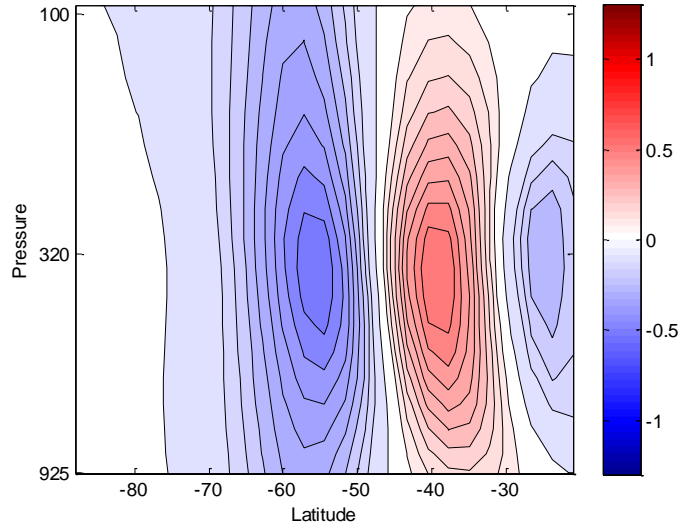


# Complex conjugate eigenvalues/eigenvectors imply propagation

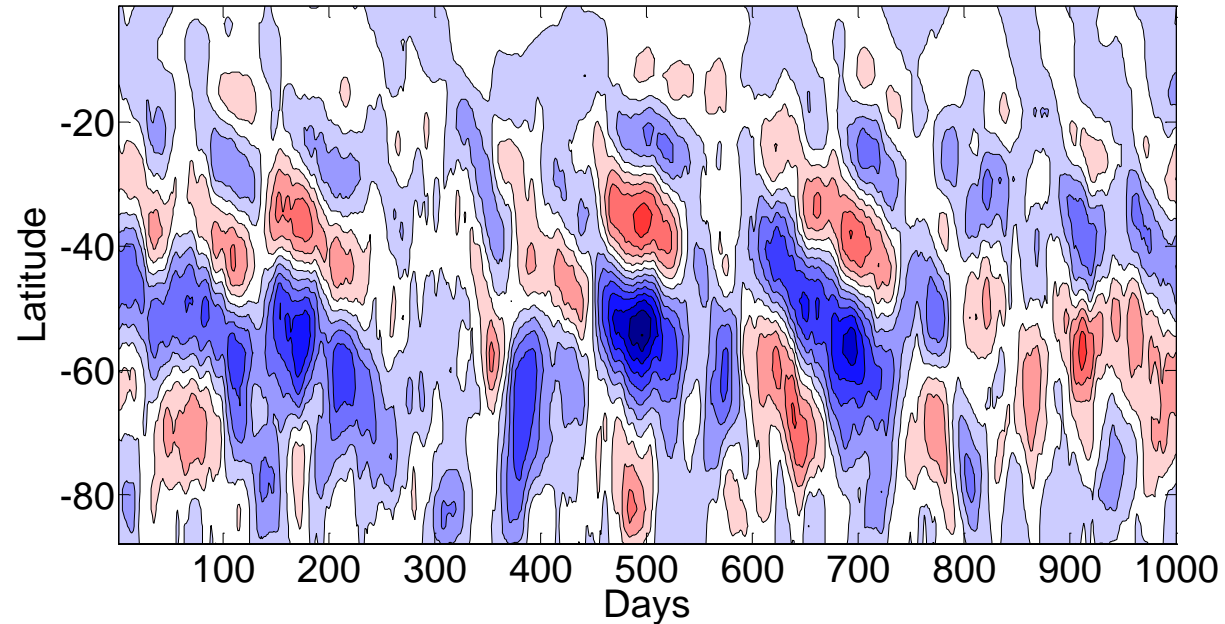
2 leading EOFs explain 40% and 29% of variance, with  $\tau = 19$  and 13 days

POP, (Re/Im)  $\Gamma=0.3574\pm 0.4818i$

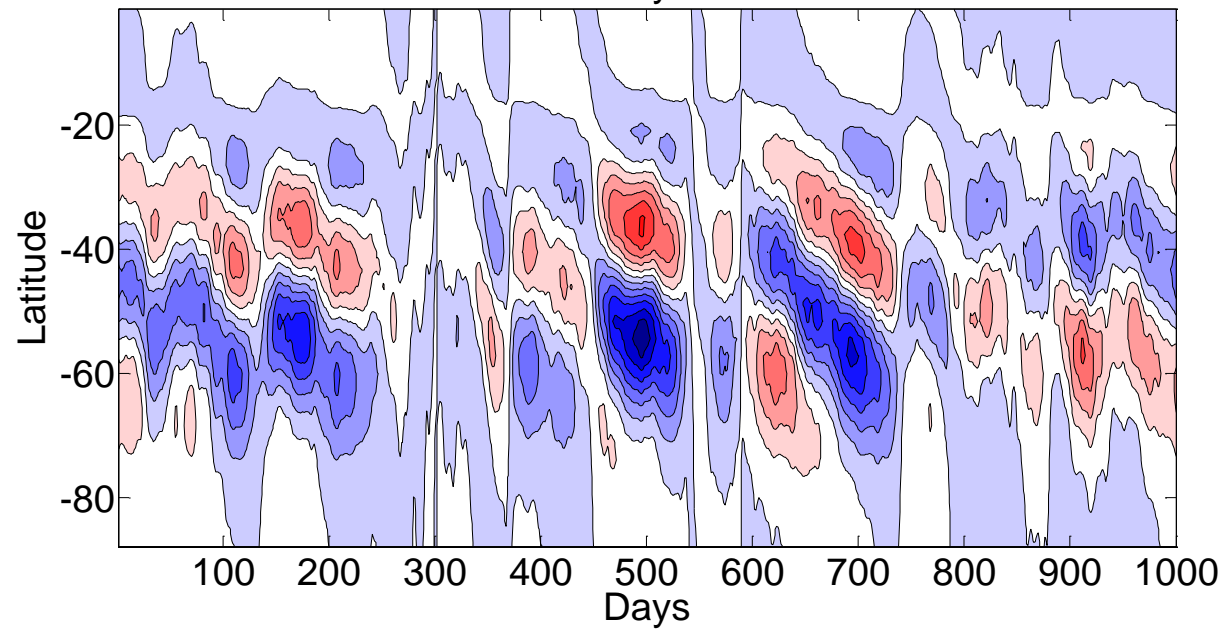
Gives a decay time of 39 days, and a period of 134.8 days (recovers the lag at which cross-correlation maximizes)



# Poleward propagation of wind anomalies



[u] anomalies integrated through the troposphere show consistent poleward propagation.. (cf. Feldstein 1998)

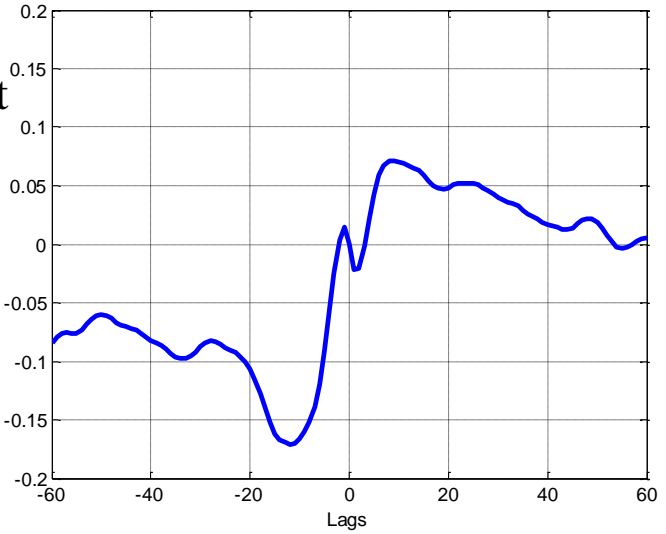


and are well captured by the first 2 EOFs

# ERA-I consistent with this picture

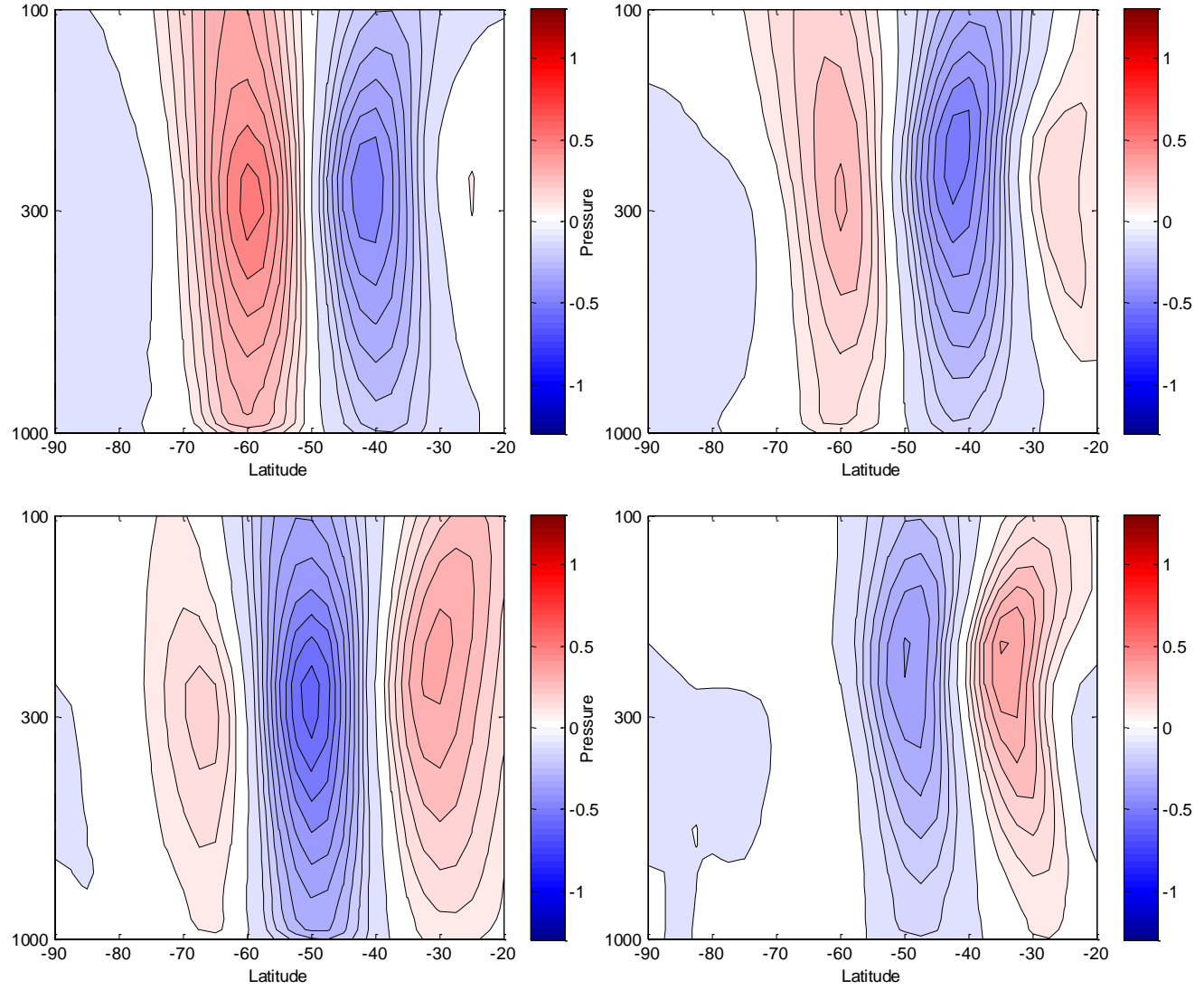
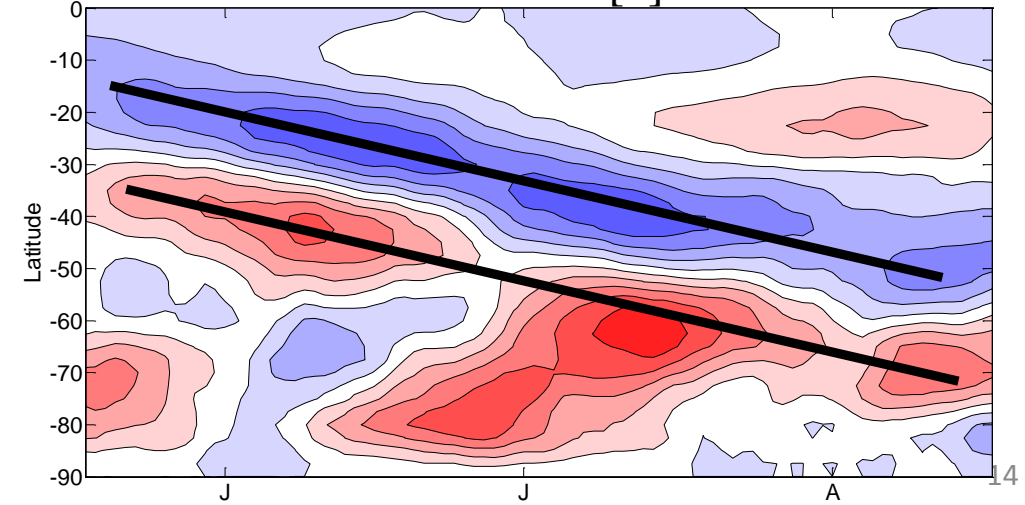
2 leading EOFs explain 37% and 19% of variance, with  $\tau = 11$  and 8 days

PCs are not independent



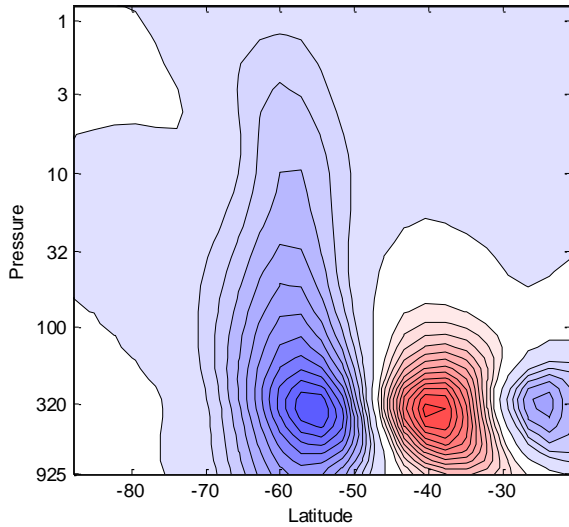
Poleward propagation of [u] anomalies

2004

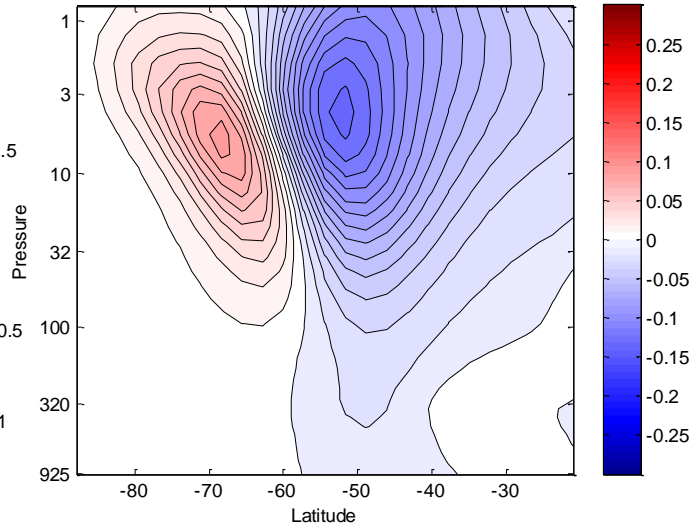


# POPs are independent of pressure weighting

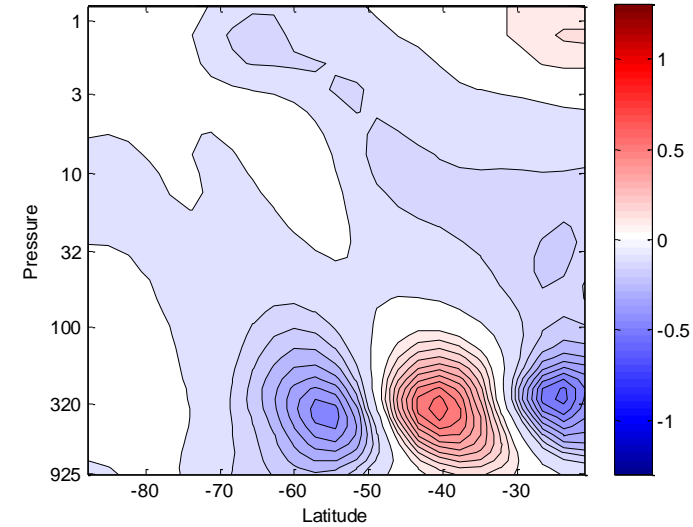
EOF1 weighted



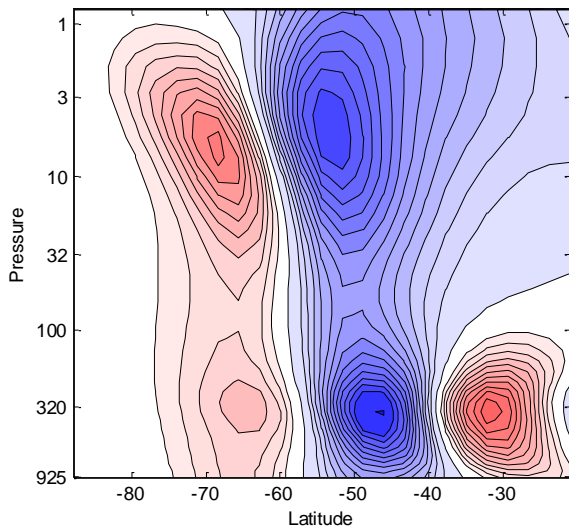
EOF1 unweighted



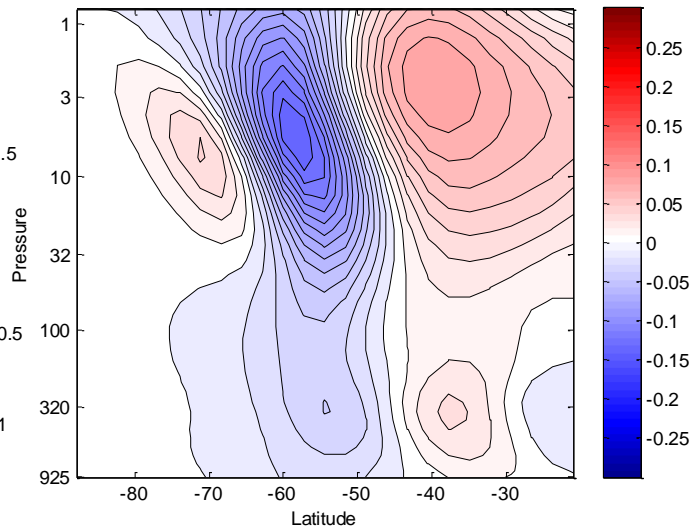
Re (POP) weighted/unweighted



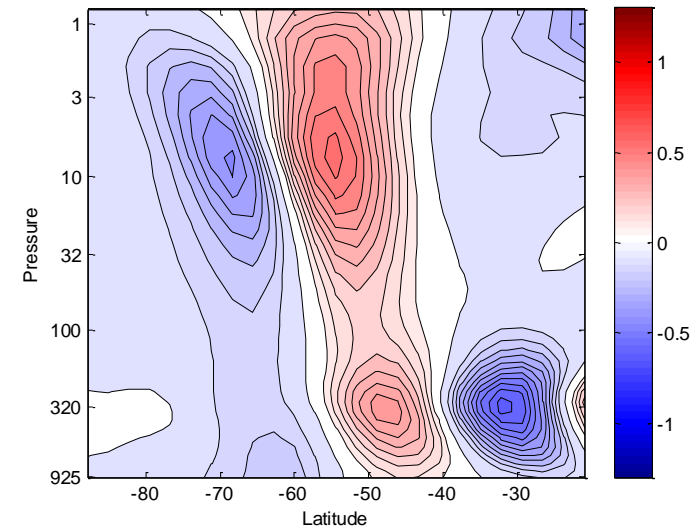
EOF2 weighted



EOF2 unweighted



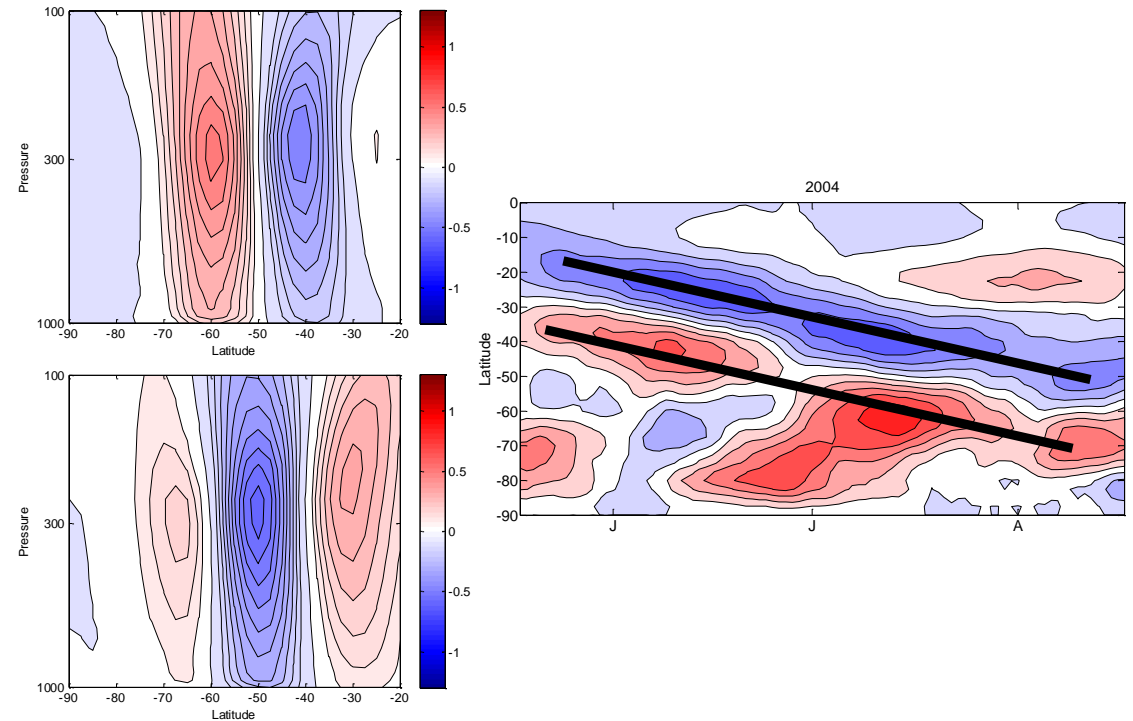
Im (POP) weighted/unweighted



# Summary – propagating annular modes

- 2 leading EOFs form one coupled, propagating mode, describing systematic latitudinal jet migration

- Models need to get POPs **and** their timescales correct



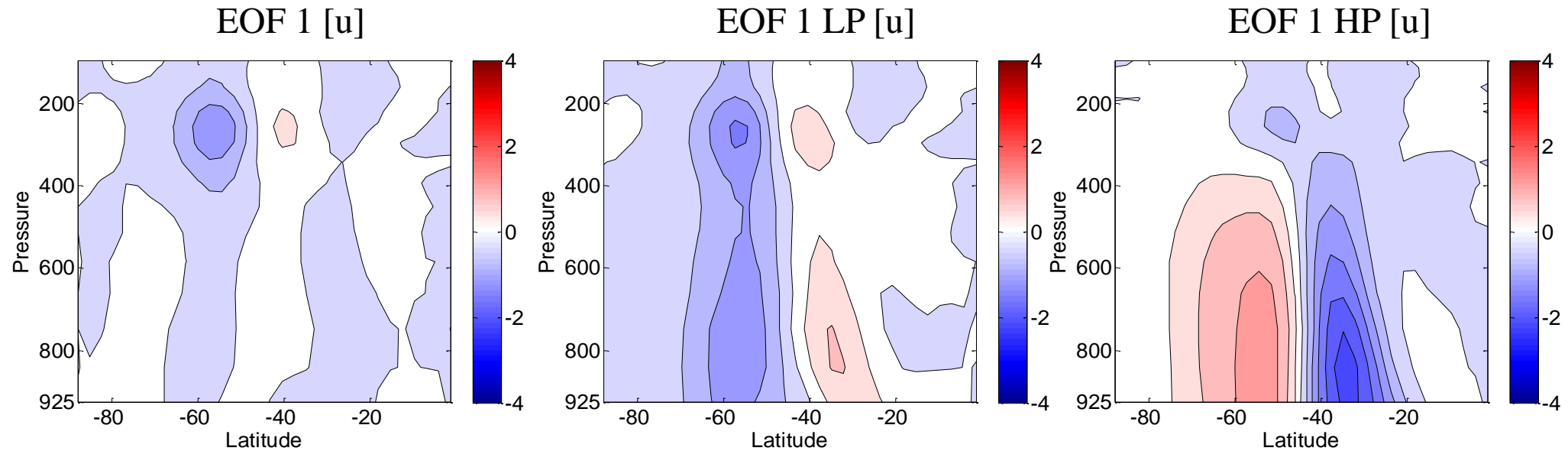
- Future work...how does external forcing project onto this mode? More details in a paper in preparation, to be submitted soon



# Frequency-dependent barotropic/baroclinic modes

Regressions of heat flux on..

(cf. Thompson and Woodworth, Thompson and Barnes)



Sheshadri and Plumb 2016b (in review at J. Atmos. Sci.)