# Timescales of eddy activity in the Southern Ocean

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## Southern Ocean Momentum Balance



Existing Theoretical Framework (Gill et. al. 1974, Marshall & Speer, 2012; Hallberg & Gnanadesikan, 2006; Abernathey & Cessi, 2014): Competition between wind driven upwelling and baroclinic eddies determines mean isopycnal slope, ACC transport, and MOC

Meredith & Hogg (2006, GRL)



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$$\frac{dP(t)}{dt} = f(t) - cP(t) \qquad \begin{tabular}{c}{\mbox{GM type}\\ \mbox{closure}\end{tabular}} \\ \frac{dK(t)}{dt} = cP(t) - rK(t) \qquad \begin{tabular}{c}{\mbox{GM type}\\ \mbox{closure}\end{tabular}} \\ \end{tabular}$$

$$f = \hat{f}e^{i\omega t} = \left|\hat{f}\right|e^{i\omega t + \phi_f}$$
$$P = \hat{P}e^{i\omega t} = \left|\hat{P}\right|e^{i\omega t + \phi_P}$$
$$K = \hat{K}e^{i\omega t} = \left|\hat{K}\right|e^{i\omega t + \phi_K}$$

$$c \sim \frac{K_{GM}}{L_y^2} \approx 10^-9 \ s^{-1}$$
 (~3 years !)







## Simple model

#### w/o Eddy feedback

$$\frac{dP(t)}{dt} = f(t) - cP(t)$$
$$\frac{dK(t)}{dt} = cP(t) - rK(t)$$

## Simple model w/o Eddy feedback with Eddy feedback

$$\frac{dP(t)}{dt} = f(t) - cP(t)$$
$$\frac{dK(t)}{dt} = cP(t) - rK(t)$$

based on mixing length arguments

$$\begin{array}{l} \frac{dP}{dt} = f - kPK^{\alpha} \\ \frac{dK}{dt} = kPK^{\alpha} - r_{1}K^{\beta} \\ & \text{general} \\ & \text{bottom drag} \end{array}$$

linearize and solve for

$$\frac{dP'}{dt} = f' - c_1 P' - c_2 K'$$
$$\frac{dK'}{dt} = c_1 P' + c_2 K' - rK'$$

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linearize and solve for

$$\frac{dP'}{dt} = f' - c_1 P' - c_2 K'$$

$$\frac{dK'}{dt} \neq c_1 P' + c_2 K' - rK'$$





based on mixing length arguments  $\frac{dP}{dt} = f - kPK^{\alpha}$  $\frac{dK}{dt} = kPK^{\alpha} - r_1K^{\beta}$ general bottom drag

linearize and solve for

$$\frac{\frac{dP'}{dt} = f' - c_1 P' - c_2 K'}{\frac{dK'}{dt} \neq c_1 P' + c_2 K' - rK'}$$

## Isopycnal GOLD model:

(Hallberg & Gnanadesikan, 2001, 2006; Howard et al. 2015)



#### Sinha & Abernathey (submitted to JPO)

- Reduced gravity model
- 4 km horizontal resolution
- Three isopycnal layers
- Wind forcing only

#### **Seven experiments**

- Steady sinusoidal wind jet (0.2 N/m^2)
- plus oscillations +/- (0.1 N/ m^2), 0.25, 0.5,1, 2, 4, 8 year periods

#### Diagnostics

- EKE
- APE
- Wind Energy input

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## Spectral Analysis











## Compare Analytic & Numerical Model



Sinha & Abernathey (submitted to JPO)

## Summary

- Two limits : Fast vs Slow Transient response to changing winds
- Analytical model: Energy Budget wind power, APE, EKE
  - with and without eddy feedback
- smooth transfer function, complex phase and amplitude response to changing winds: Regime shift
- Numerical simulations with idealized model
- mechanistic description of the eddy equilibration process with purely dynamic forcing

# Discussions

- Eddy generation and dissipation non-local in time
- Eddy memory effect Time dependent eddy parameterization
- Used in conjunction with multiple timescale response to thermodynamic forcing (Ferreira et al 2014)(sea ice, ozone depletion etc.) - more complete theory for SO response, baroclinic eddy equilibration

Thank you.









Appendix

 Meredith and Hogg (2006) - EKE peaks around 2-3 years after SAM : Wind Energy stored as PE slowly transferred to EKE



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165

165

170

170

175

175

180









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 But there is more than one timescale !

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$$W + B_{ML} = \frac{d}{dt}(PE) + \frac{d}{dt}(KE) + D + V$$

160

160

165

165

time (year)

170

170

175

175

180





But there is more than one timescale !

• Treguier et. al. (2010): response to increase in SAM in eddy permitting model

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#### Simple model with Eddy Feedback



#### Power Spectra of Winds

#### NCAR Reanalysis

#### **ERA** Interim



### Spectral Amplitude Response



#### Transport Spectra





